A nonconservative hyperbolic depth-averaged model for granular-fluid mixtures: application to debris flows and landslides.

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- Software packages: numerical methods summary
- A few examples
- Debris flows: physics/modeling
- Debris flows: our mathematical model
- Debris flows: some numerical issues

- Open-source software package for general hyperbolic problems:
- High-resolution (2nd-order) Godunov methods;
- Finite volume discretization on logically rectangular grids;
- Grid mappings for various geometries;
- Block-structured adaptive mesh refinement;
- Wave-propagation (possibly non-conservative) update from Riemann Solutions;
- Available @ www.clawpack.org.

## The GeoClaw Software Project

- Subset of Clawpack for hyperbolic problems in geophysics:
  - free-surface flows (tsunamis, flooding, landslides etc.)
  - volcanic plume dynamics, ash clouds, pyroclastic flows etc.
  - seismic wave propagation
  - subsurface flow problems (mixed-type equations)
- Adaptive mesh refinement (AMR) schemes tailored to free-surface flows with inundation;
- Positive-depth preserving Riemann solvers;
- General inundation Riemann solver ability;
- Well-balanced Riemann solvers (to varying degree) for depth-averaged flows;
- Various user tools for geophysical data:
  - e.g., dynamic processing of unaligned topography DEMs;

## GeoClaw Example: Honshu-Tohoku 2011 tsunami

#### GeoClaw Example: inundation of Hilo, Hawaii

Using 5 levels of refinement with ratios 8, 4, 16, 32. Resolution  $\approx 160$  km on Level 1 and  $\approx 10$ m on Level 5. Total refinement factor:  $2^{14} = 16,384$  in each direction.



George Debris-Flow Model

## GeoClaw Example: inundation of Hilo, Hawaii



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Flows characterized by a variable granular-fluid mixture:

- tsunami propagation and inundation;
- hurricane generated storm-surge inundation;
- overland/fluvial flooding, dam and levee breaches etc.;
- sediment erosion, deposition and transport;
- landslides, mudslides, lahars and debris flows;
- dry granular (or snow) avalanches.







Common mathematical features and computational challenges:

- flow is shallow relative to length-scales;
- often modeled with nonconservative hyperbolic systems;
- flow moves over complex topography (singular sources);
- domain is of varying bounded extent (wet/dry problem);
- dynamics are a small perturbation to a steady state;
- feature evolving multiple spatial scales.







## Hazardous Earth-Surface Flows

Depth-averaged models: flow between a fixed bottom b(x, y) with a free surface  $\eta(x, y, t)$ :



Shallow assumption and B.C.s give two-dimensional systems:

$$q_t + A(q)q_x + B(q)q_y = \psi(q),$$
$$q = (h, hu, hv, \dots,)^T$$

Modeling fluidized-granular flow requires a stress model





Indonesian Debris Flow Movie Ritigraben Switzerland

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A true Debris flow is best described as a two or three-phase flow

- two-phase models treat the fluid and suspended fines as one phase
- shallow assumption usually holds (depth-averaging works)
- fluid-fluid stress (pressure, viscosity)
- solid-solid stress (Coulomb-Mohr friction, collisional)
- solid-fluid stress (pressure, drag)

Common categories of debris-flow models

- 3D models:
  - mixture models: single continuum rheology (usually non-Newtonian)
  - two-phase models: (usually too expensive for large-scale problems.
- Depth-averaged models:
  - Savage-Hutter models: dry-granular flows (Mohr-Coulomb friction).
  - Depth-averaged mixture models: single continuum rheology.
  - Depth-averaged multi-phase models: complex stress models for each phase
  - Quasi-two phase models: only most important contribution from fluid phase are modeled.

Our debris-flow model (physical fidelity  $\Leftarrow \Rightarrow$  model tractability)

- quasi-two-phase:
  - mass (depth), momentum and volume fractions for solid phase are retained.
  - specific discharge is neglected. (fluid drag on solid is ignored).
  - volume-fractions are retained.
  - fluid pressure is evolved and is coupled with solid-volume fraction and momentum.
- stress model:
  - Coulomb friction for solid-solid stress: (mediated by pore-fluid pressure)
  - fluid effect on solid phase is through pressure
  - pore-fluid pressure is strongly coupled to solid-volume fraction

1D model on a fixed incline:



System for 
$$q = (h, hu, hm, p)^T$$
.  

$$h_t + (hu)_x = D \frac{(\rho - \rho_f)}{\rho},$$

$$(hu)_t + (hu^2 + \frac{1}{2}\kappa g^y h^2)_x + \frac{(1 - \kappa)}{\rho} p_x = g^x h + uD \frac{(\rho - \rho_f)}{\rho} - \frac{\tau}{\rho},$$

$$(hm)_t + (hum)_x = -Dm \frac{\rho_f}{\rho},$$

$$p_t - \gamma \rho g^y uh_x + \gamma \rho g^y (hu)_x + up_x - = \zeta D - \frac{3}{h\alpha(1 + \kappa)} u \tan(\psi)$$

where,

$$D = \frac{k}{h\mu} (\rho_f g^y h - p).$$

System for 
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These equations are a hyperbolic system of the form:

$$q_t + f(q)_x + W(q)q_x = \psi(q, x)$$

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Eigenstructure of  $A(q), q = (h, hu, hm, p)^T$ :  
$$r_{1,2,3,4} = \begin{bmatrix} 1 \\ u - \sqrt{g^y h \epsilon} \\ m \\ \gamma \rho g^y \end{bmatrix}, \begin{bmatrix} 0 \\ 0 \\ 1 \\ 0 \\ 1 \\ 0 \end{bmatrix}, \begin{bmatrix} (\kappa - 1) \\ u(\kappa - 1) \\ 0 \\ \kappa \rho g^y \end{bmatrix}, \begin{bmatrix} 1 \\ u + \sqrt{g^y h \epsilon} \\ m \\ \gamma \rho g^y \end{bmatrix}$$

)

,

 $\lambda_{1,2,3,4} = u - \sqrt{g^y h \epsilon}, u, u, u + \sqrt{g^y h \epsilon}$ 

Always a full set of eigenvectors:

$$|r_{1,2,3,4}| = 2\rho g^y \sqrt{g^x h\epsilon}$$

where,

$$\epsilon = \gamma + (1 - \gamma)\kappa, \quad 3/4 < \gamma < 1 \Rightarrow \epsilon > 0$$

Mobilization:  $q_t + f(q)_x + W(q)q_x = \psi(q, x)$ 

- failure occurs when the driving forces slightly exceed the shear strength at any one point: a small perturbation to a balanced steady state:  $f(q)_x + W(q)q_x \approx \psi(q,x)$
- failure: an equilibrium is perturbed...what happens next?
  - **1** rapid temporary instability: (shear contraction  $\rightarrow$  increased pore pressure  $\rightarrow$  decreased shear strength  $\rightarrow$  landslide.)
  - 2 quick stabilization: (shear dilation  $\rightarrow$  decreased pore pressure  $\rightarrow$  shear strength reestablished  $\rightarrow$  localized slump.)

3 anything intermediate...e.g., stick-slip.

 modeling the outcome at least requires well-balanced methods if it can be done at all

Well balancing:  $q_t + f(q)_x + W(q)q_x = \psi(q, x)$ 

- f-wave approach:  $\Delta f + W(\bar{Q})\Delta Q \Psi = \sum \beta_p r_p$
- steady-state wave approach:

• 
$$q_t + f(q)_x + W(q)q_x = \psi(q, x)$$

• 
$$q_t + A(q)q_x = \psi_1(q, x) + \psi_2(q, x)$$

• 
$$\tilde{q}_t + A(\tilde{q})\tilde{q}_x = 0; \tilde{q}_t = \psi_2(q, x)$$

• 
$$\tilde{R}(\tilde{Q})^{-1}\Delta\tilde{Q} = \tilde{\beta}$$

• steady-state solution:  $\tilde{A}(\tilde{q})\tilde{q}_x = 0;$ 

• 
$$\tilde{R}(\tilde{Q})^{-1}(\tilde{Q}_r - \tilde{Q}_l) = (0, \dots, \tilde{\beta}_0, \dots, 0)^T.$$

• 
$$Q_r - Q_l = \beta_0 \tilde{r}_0(Q).$$

• 
$$A(\tilde{Q})(\tilde{Q}_r - \tilde{Q}_l) = \lambda_0(\tilde{Q}_r - \tilde{Q}_l) = 0$$

- Balanced steady-states arise:  $A(q)q_x \approx \psi(q, x)$ .
- Role of source term must be treated carefully in the Riemann problem.



- We solve augmented homogeneous system:  $W_t + \mathcal{A}(W)W_x = 0.$
- steady states: stationary steady state wave only:  $\mathcal{A}(\mathcal{W})\mathcal{W}_x=0$



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motionless steady state at the shoreline: no waves.







































## USGS experimental debris-flow flume



Flow Dynamics Movie Mobilization Movie Flow Dynamics and Segregation

# **Testing mobilization**

## **Testing flow dynamics**

Model test:

## **Testing flow dynamics**



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**Debris-Flow Model** 

## **Testing flow dynamics**





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**Debris-Flow Model** 

- Future Directions
  - incorporation of particle size segregation
  - entrainment
  - arbitrary topography
  - vertical acceleration corrections?
- Thank you for your attention.